## Astronomy Summary Knowledge Organiser - Chapter 11 (Topic 4) Time and the Earth-Moon-Sun cycles (i)

The SIDEREAL DAY is measured
with respect to the STARS.

It lasts 23 hours 56 minutes and 4 seconds. This is the time between a star completing two successive passages across the observers meridian (the line connecting the due north and south points on the observers horizon, that also passes through their zenith). To cause this to happen the Earth must rotate $360^{\circ}$.
The length of Earth's sidereal day (its sidereal period) can be found by taking long exposure photographs' of stars moving around the North Celestial Pole(see below) and measuring the mean angle that the stars appear to have moved (angle subtended) during that exposure time. The formula below the image then allows you to calculate the length of the sidereal day.


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The SOLAR DAY (or SYNODIC DAY) is measured using the movement of the MEAN SUN across the sky. It lasts 24 hours.
During each solar day, the motion of the Earth as it continues its orbit around the Sun means the Earth actually rotates $361^{\circ}$ with respect to the background stars, before the Sun returns to the same position in the sky.


The four seasons and the annually changing rising and setting times of the Sun are caused by a combination of the Earth's orbit of the Sun and its axis of rotation being tilted by $23.5^{\circ}$.
At the SPRING \& AUTUMNAL EQUINOX the Earth's axis is tilted at $90^{\circ}$ relative to the Sun and so days $\&$ nights are roughly equal in length. On this date the Sun will be directly overhead at midday if you are stood on the Equator

| Spring equinox $=$ March $21^{\text {st }}$ | Summer Solstice $=$ June $21^{\text {st }}$ |
| :--- | :--- |
| Autumnal equinox $=$ September | $23^{\text {rd }}$ |
|  | Winter Solstice $=$ Dec $21^{\text {st }}$ |

On the SUMMER SOLSTICE the Earth's axis (Northern Hemisphere) is tilted directly towards the Sun and so this is the longest day of the year and so has the shortest night. On this date the Sun will be directly overhead at midday if you are stood at a latitude of $23.5^{\circ}$ North (on the Tropic of Cancer).

For the WINTER SOLSTICE the opposite is true and the Sun can be seen directly overhead on the Tropic of Capricorn (23.5 South).

The Moon completes one orbit of Earth in 27.3 days, this is called a SIDEREAL MONTH. However, due to the continuing motion of the Earth around the Sun during that time, it takes 29.5 days to get from New Moon to New Moon, this is called the SOLAR SYNODIC MONTH or lunar phase cycle. Because the Earth completes $27^{\circ}$ of its orbit of the Sun whilst the Moon completes one orbit of the Earth, the Moon has to continue its motion for an extra 2.2 days before it is again directly aligned with the Sun and Earth.


1 New moon 2 Waxing crescent 3 First quarter half moon
4 Waxing gibbous 5 Full moon 6 Waning gibbous
7 Third quarter half moon 8 Waning crescent

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The traditional SUNDIAL (shown to the right) has a triangular metallic GNOMON that points to the North Celestial Pole. The gnomon casts a shadow onto a graduated scale on the flat surface below from which the time can be read. Sundials tell us the APPARENT SOLAR TIME (AST). Clocks tick by the same amount every second of every day of the year but sadly the motion of the Sun across the sky is not regular, sometimes it appears to move fast and some times more slowly. Due to this the AST read of a sundial usually needs a CORRECTION to change it to a MEAN SOLAR TIME (MST), based on a MEAN SUN that moves regularly across the sky!
The correction applied to the AST to change it to the MST that we use in our 24 hour clocks is called the EQUATION of TIME (EOT).
The value of the EOT can be negative or positive and changes throughout the year, it is shown on the graph below.


The following formula allows you to calculate the EOT, AST or MST;
$E O T=A S T$ - MST
Rearranging,
MST = AST - EOT



The image above is an ANALEMMA that shows the variation of the EOT along with both the Sun's declination \& month

If the EOT is a positive value it is because the real Sun is fast (and so AST will be ahead of MST). A negative EOT means a slow real Sun.
If a sundial reads 11.20 on a day the EOT is $\mathbf{- 1 3}$ mins, then the MST is 11.33
If you look in the ANALEMMA below (the 'wobbly snowman') it must be the month of February and the Sun will be low in the sky!
The annual variation in the EOT is CAUSED BY;
The EARTH'S ELLIPTICAL ORBIT - Earth travels around the Sun at different speeds in its orbit, fastest at perihelion and slowest at aphelion!
The TILT of the EARTH'S AXIS - in the months close to the summer \& winter solstice the real Sun moves across the sky (East to West) faster than in the months closest to the Spring \& Autumnal Equinox. At times close to the equinoxes a large component of the real Sun's apparent motion is northwards or southwards and therefore the Sun's east to west motion laas behind the mean Sun.


A SHADOW STICK is a thin vertical stick that casts a shadow that can be measured in length, at different times during the day, to find LOCAL NOON \& LONGITUDE As the Sun rises in the east \& sets in the west, it will be low in the sky and so the

## shadows will be long



When the Sun culminates at local noon, it will be high in the sky and so the shadow cast will be at its shortest! This is shown in the graph above.
The difference between the time of local noon shown on the graph (lowest point of the line of best fit plotted)(AST) and the time read on a watch, phone or clock (MST) is due to the EOT and/or the observers longitude.

## Astronomy Summary Knowledge Organiser - Chapter 11 (Topic 4) Time and the Earth-Moon-Sun cycles (iii)

If data from a shadow stick experiment shows that it is local noon (AST $=12.00$ ) on a day when the EOT is -8 mins we can use the EOT formula to calculate the Local Mean Time (shown below):
EOT = AST - MST rearranged:
MST $=A S T-E O T$
MST $=12.00-(-8)$
MST $=12.00+8$
MST $=12.08$
This calculation suggests the Local Mean Time to be 12.08 but the MST (as shown on civil watches) might in fact be 12.04 (4 minutes difference). This final difference between the calculated Local Mean Time and the MST shown on watches must be due to the longitude of the observer. If the Local Mean Time calculated by the shadow stick is 4 minutes ahead of the MST read on a watch or phone, the observer must be $1^{\circ}$ East of the Greenwich Meridian. So, long. $=1^{\circ}$ East.

$$
1^{\circ}=4 \mathrm{~min} \quad 5^{\circ}=20 \mathrm{~min} \quad 15^{\circ}=1 \text { hour }
$$

Any observations that produce a Local Mean Time that is EARLIER than the MST(Greenwich Mean Time) shown on a watch or phone prove that the observers longitude is WEST of the Greenwich Meridian. e

Every 4 mins of time is equivalent to one degree of longitude. On the image to the right, at any location East of the Primel Greenwich Meridian, the LMT calculated will be later than GMT.


## The LONGITUDE at SEA PROBLEM

In the $17^{\text {th }}$ Century there was a rapid expansion in the number of ships going off to sea to explore new territories, to wage war or to trade gold, silk or spices. Captains could easily determine their ships LATITUDE by using the altitude of either Polaris at night or the Sun at local noon. However, calculating a ships LONGITUDE whilst out on the ocean was a huge problem that caused many shipwrecks and fatalities!

In theory, finding longitude at sea was a simple process - use the position of the Sun to find the Local Mean Time on the ship and then compare it to the Local Mean Time in the port from where the ship sailed - the difference in time (after taking the EOT into account) could then be converted into degrees of longitude, east or west of the port. However, knowing the LMT back at the port after a long journey required a clock that kept accurate time, even on rough seas, this was beyond the clockmakers of the time.

Eventually, after many attempts failed, in 1761, the horologist JOHN HARRISON solved the longitude problem when he invented and manufactured a MARINE CHRONOMETER (pictured right) that was capable of keeping accurate 'home-port' time, even aboard ships that were sailing in conditions of raging seas, extreme temperatures, high humidity and salinity!


Before rail travel became common in Britain in the mid 1800's people relied on Local Mean Time in all parts of life. But since LMT differed from city to city due to their different longitudes this caused huge problems in such things as creating train timetables. Finally, in 1880 the Government standardised time for the whole country with the local mean solar time at Greenwich in London, it was called GREENWICH MEAN TIME (GMT).
Years later, other countries followed the British example and so the Earth was eventually split into one hour time zones ( $15^{\circ}$ of longitude wide).



[^0]:    length of sidereal day $=\frac{\text { exposure time } \times 360^{\circ}}{\text { mean angle subtended by arcs }}$

